

## Mass Wasting

**Mass wasting:** Mass wasting is a term used to describe the movement of geological material (rocks, soils, sediments) under the effects of gravity but frequently affected by water and water content as in submarine environments and mud slides. Types of mass wasting include **creep, slides, flow, topples, falls,** and **sinkhole** each with its own characteristic features, and taking place over timescales from seconds to years. It proceeds at variable rates of speed and is largely dependent on the **water saturation levels** and the **steepness** of the terrain. A destructive, rapid mass-wasting event is called a **landslide**; if the movement is slow enough that it cannot be seen in motion, it is called **creep**.

**Three kinds of movement** are generally recognized: **flow, slip,** and **fall**. A mass-wasting event is called a **flow** if the mass moves down slope like a viscous fluid. If the mass moves as a solid unit along a surface or plane, it is called a **slip**. A **slip** that moves along a surface parallel to the slope is called a **slide**. If the movement occurs along a curved surface where the downward movement of the upper part of the mass leaves a steep scarp (cliff) and the bottom part is pushed outward along a more horizontal plane, it is called a **slump**. Earth material that free-falls from a steep face or cliff is termed a **fall**.

### Mass-Wasting Controls

Varieties of conditions affect the development of mass wasting in a particular area. 1- Steep slopes, 2- widely varying altitude ranges (relief), 3- the thickness of the loose earth material, 4- planes of weakness parallel to the slopes, 5- frequent freezing and thawing, 6- high water content in the earth material, 7- dry conditions with occasional heavy rainfall, 8- and sparse vegetation are the factors that contribute to the unstable conditions that result in mass wasting. Movements can be triggered by the motion of earth quakes or too much weight added to the upper part of a slope, such as snowpack.

## FACTORS THAT CONTROL MASS WASTING

### 1- Steepness of the slope

Obviously, the steepness of a slope is a factor in mass wasting. If frost wedging dislodges a rock from a steep cliff, the rock tumbles to the valley below. However, a similar rock is less likely to roll down a gentle hill side.

### 2-Type of rock and orientation of rock layers

If sedimentary rock layers dip in the same direction as a slope, the upper layers may slide over the lower ones. Imagine a hill underlain by shale, sandstone, and limestone oriented so that their bedding lies parallel to the slope, as shown in Figure 6–1a. If the base of the hill is undercut (Fig. 6–1b), the upper layers may slide over the weak shale. In contrast, if the rock layers dip at an angle to the hill side, the slope may be stable even if it is undercut (Figs. 6–1c and 6–1d).

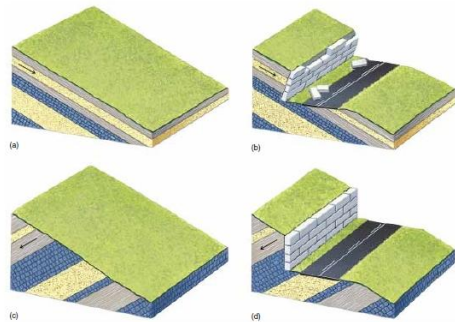


Figure 6–1 (a) Sedimentary rock layers dip parallel to this slope. (b) If a road cut undermines the slope, the dipping rock provides a good sliding surface, and the slope may fail. (c) Sedimentary rock layers dip at an angle to this slope. (d) The slope may remain stable even if it is undermined.

Several processes can undercut a slope. A stream or ocean waves can erode its base. Road cuts and other types of excavation can also destabilize it. Therefore, a geologist or engineer must consider not only a slope's stability before construction, but how the project might alter its stability.

**3- Angle of repose:** The angle of repose is the steepest angle at which loose material will remain in place. It is largely dependent on the size, shape, and roughness of the particles. The angle varies from about 25 degrees to about

40 degrees. If the angle is exceeded by additional sedimentation or tilting, a slide or disturbance will result.

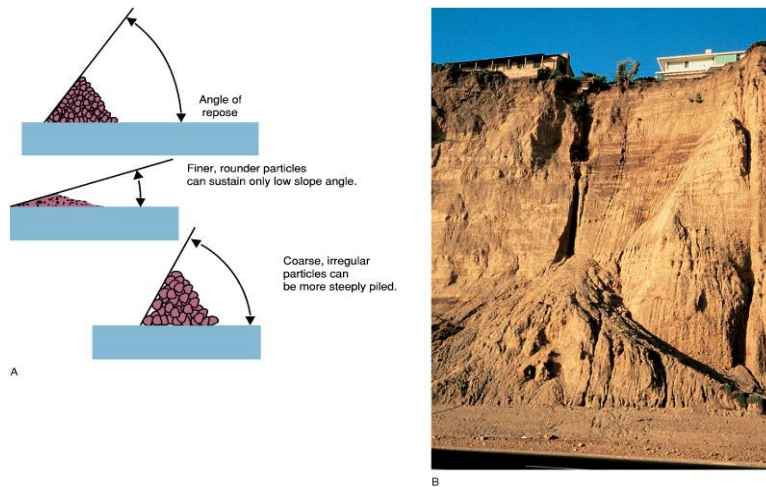


Figure 6–2 (A), shows the variation of the angle of repose due the size , shape, and roughness, B, photo of angle of repose.

#### 4-Water and vegetation

To understand how **water** affects slope stability, think of a sand castle. Even a novice sand-castle builder knows that the sand must be moistened to build steep walls and towers (Fig. 6–3). But too much water causes the walls to collapse. Small amounts of water bind sand grains together because the electrical charges of water molecules attract the grains. However, excess water lubricates the sand and adds weight to a slope. When some soils become water saturated, they flow downslope, just as the sand castle collapses. In addition, if water collects on impermeable clay or shale, it may provide a weak, slippery layer so that overlying rock or soil can move easily.



Figure 6–3 The angle of repose depends on both the type of material and its water content. Dry sand forms low mounds, but if you moisten the sand, you can build steep, delicate towers with it.

**Roots** hold soil together and plants absorb water; therefore, a highly vegetated slope is more stable than a similar bare one. Many forested slopes that were stable for centuries slide when the trees were removed during logging, agriculture, or construction.

**Vegetation** is sparse because of summer drought and wildfires. When winter rains fall, bare hillsides often become saturated and slide. Mass wasting occurs for similar reasons during infrequent but intense storms in deserts.

## **5-Earthquakes and volcanoes**

An earthquake may cause mass wasting by shaking an unstable slope, causing it to slide. A volcanic eruption may melt snow and ice near the top of a volcano. The water then soaks into the slope to release a landslide.

## **TYPES OF MASS WASTING**

Mass wasting can occur slowly or rapidly. In some cases, rocks fall freely down the face of a steep mountain. In other instances, rock or soil creeps down slope so slowly that the movement may be unnoticed by a casual observer.

### **1-FLOW**

Types of flow include **creep**, **debris flow**, **earthflow**, **mudflow**, and **solifluction**.

#### **Creep**

As the name implies, creep is the slow, downhill movement of rock or soil under the influence of gravity. Individual particles move independently of one another, and the slope does not move as a consolidated mass. A creeping slope typically moves at a rate of about 1 centimeter per year, although wet soil can creep more rapidly. During creep, the shallow soil layers move more rapidly than deeper material (Fig. 6–4). As a result, anything with roots or a foundation tilts downhill. Over the years, soil creep has tipped the older

monuments, but the newer ones have not yet had time to tilt. Trees have a natural tendency to grow straight upward. As a result, when soil creep tilts a growing tree, the tree develops a J-shaped curve in its trunk called pistolbutt (Fig. 6-4a). If you ever contemplate buying hillside land for a home site, examine the trees. If they have pistol-butt bases, the slope is probably creeping, and creeping soil may tear a building apart.

### Creep interpretation

Creep can also result from freeze–thaw cycles in the spring and fall in temperate regions. Recall that water expands when it freezes. When damp soil freezes, expansion pushes it outward at a right angle to the slope. However, when the Sun melts the frost, the particles fall vertically downward, as shown in Figure 6-4. This movement creates a net downslope displacement. The displacement in a single cycle is small, but the soil may freeze and thaw once a day for a few months, leading to a total movement of a centimeter or more every year.

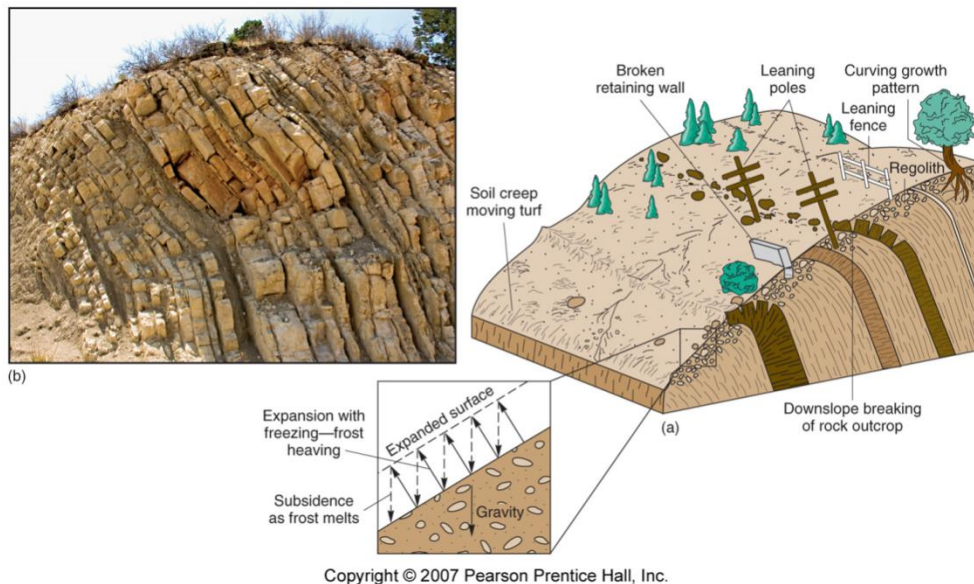


Figure 6–4, sketch of inclination of many objects like trees, poles, fences, retaining walls and layers due to creep, in lower part the interpretation of creep. b, Photo shows bending of layers.

## Debris Flows, Mudflows, and Earthflows

Different types of flows are characterized by the **sizes** of the solid particles. A **debris flow** consists of a mixture of **clay, silt, sand, and rock fragments** in which more than **half** of the particles are larger than sand (Fig. 6–5). In contrast, mudflows and earthflows are predominantly sand and mud. Some **mudflows** have the consistency of wet concrete, and others are more fluid. Because of its high water content, a mudflow may race down a stream channel at speeds up to 100 kilometers per hour. An **earthflow** contains less water than a mudflow and is therefore less fluid (Figure 6–6).

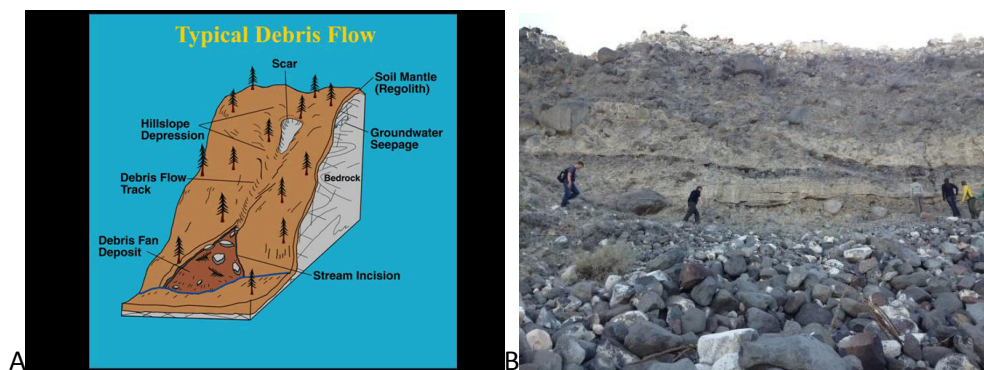


Figure 6–5, A shows sketch of debris flow, B shows an area affected by debris flow.

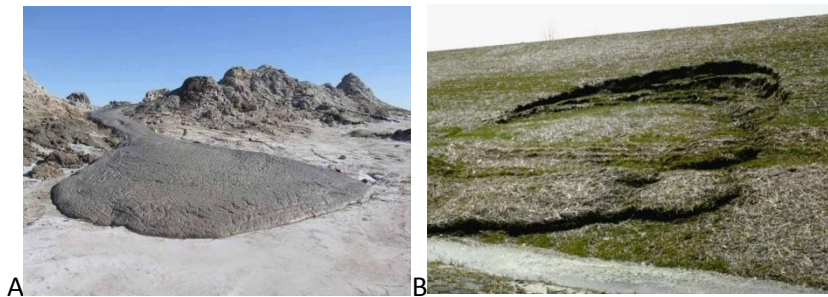


Figure 6–6, shows photo of A, mudflow, B, Earthflow.

**Solifluction** is a type of mass wasting that occurs when water-saturated soil flows downslope. It is most common in permafrost regions, where the permanent ice layer causes overlying soil to become waterlogged, although it can also occur in the absence of permafrost (Fig. 6–7). Solifluction can occur on a very gentle slope, and the soil typically flows at a rate of 0.5 to 5 centimeters per year.



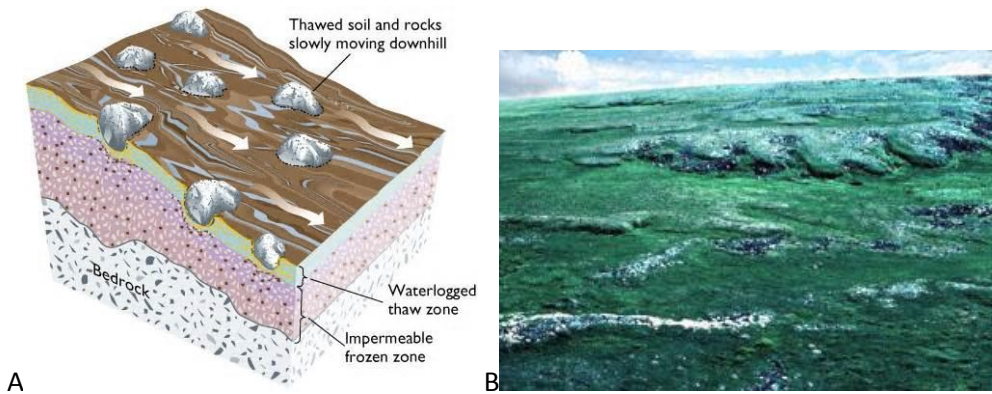


Figure 6–7, shows, A skech of solifluction, B , photo of solifluction.

## 1- SLIDE

In some cases, a large block of rock or soil, or sometimes an entire mountain side, breaks away and slides down slope as a coherent mass or as a few intact blocks. Two types of slides occur: **slump** and **rockslide**.

A **slump** occurs when blocks of material slide downhill over a gently curved fracture in rock or regolith (Fig.6-8). Trees remain rooted in the moving blocks. However, because the blocks rotate on the concave fracture, trees on the slumping blocks are tilted **backward**. Thus, you can distinguish slump from creep because slump tilts trees **uphill**, whereas creep tilts them **downhill**. At the lower end of a large slump, the blocks often pile up to form a broken, jumbled, hummocky topography. It is useful to identify slump because it often recurs in the same place or on nearby slopes. Thus, a slope that shows evidence of past slump is not a good place to build a house.

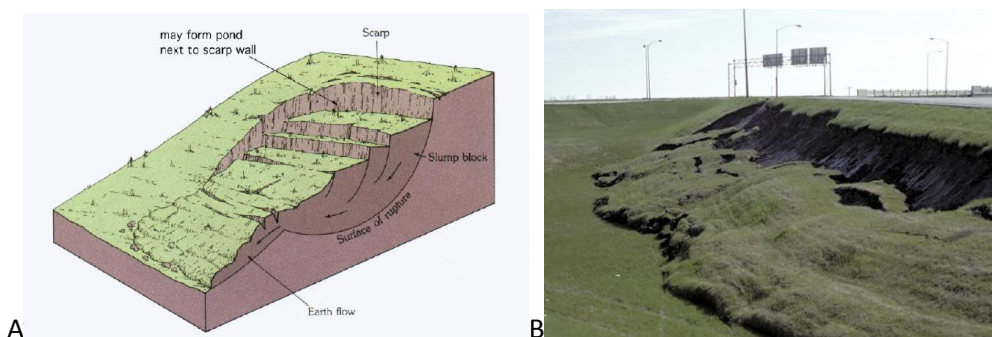


Figure 6–8, shows skech of slump, B, photo of sluped area.

During a **rockslide**, or **rock avalanche**, bedrock slides downslope over a fracture plane. Characteristically, the rock breaks up as it moves and a turbulent mass of rubble tumbles down the hillside (Fig 6-9). In a large **avalanche**, the falling debris traps and compresses air beneath and within the tumbling blocks. The compressed air reduces friction and allows some avalanches to attain speeds of 500 kilometers per hour. The same mechanism allows a snow or ice avalanche to cover a great distance at a high speed (Fig 6-10).

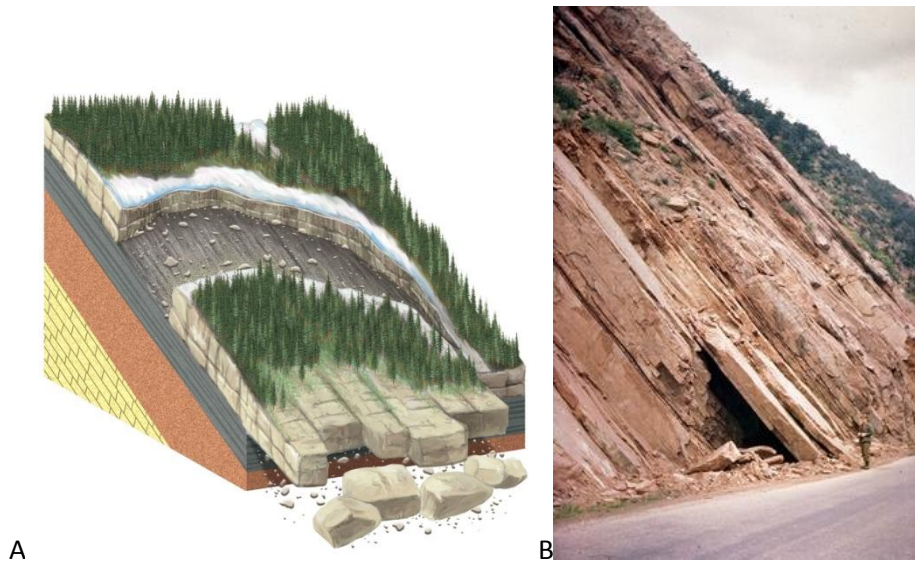


Fig 6-9, shows sketch of rock slide, B, photo of rockslide.



Fig 6-10, shows photo of rock avalanche.



## FALL

If a rock dislodges from a steep cliff, it falls rapidly under the influence of gravity. Several processes commonly detach rocks from cliffs, that when water freezes and thaws, the alternate expansion and contraction can dislodge rocks from cliffs and the growth of roots trees and cause rockfall. Rockfall also occurs when a cliff is undercut. For example, if ocean waves or a stream undercuts a cliff, rock above the waterline may tumble (Fig. 6–11).

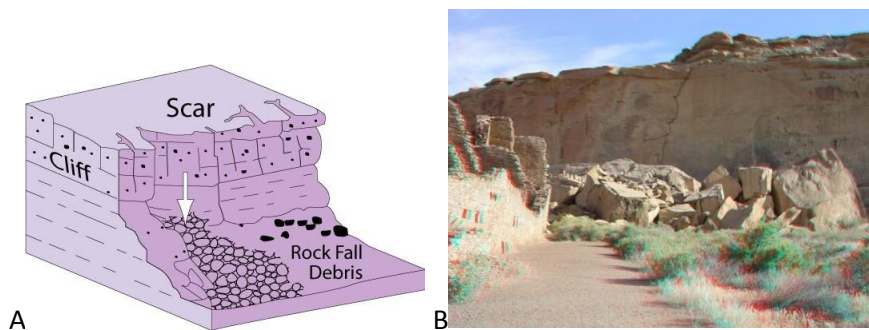


Fig 6-11, shows sketch of rockfall, B, photo of rockfalls.

**Subsidence or sinkhole**, is a depression or hole in the ground caused by some form of collapse of the surface layer, frequently causes major problems in karst terrains, where dissolution of limestone by fluid flow in the subsurface causes the creation of voids (i.e. caves). If the roof of these voids becomes too weak, it can collapse and the overlying rock and earth will fall into the space, causing subsidence at the surface. This type of subsidence can result in sinkholes which can be many hundreds of meters deep (Fig. 6–12).

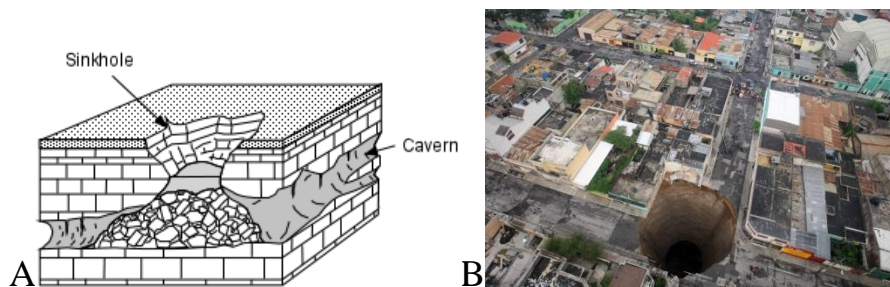


Fig. 6–12, A. shows sketch of sinkhole, B shows photo of sinkhole in Guatemala.